

Long-term effect of temperature and precipitation on radial growth in a threatened thermo-Mediterranean tree population

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Abstract

Key message Based on the first dendroclimatological analyses of the thermo-Mediterranean tree *Pyrus bourgaeana*, the positive relationship between the growth and climate (i.e., precipitation) has strengthened in recent decades.

Abstract The combined effect of climate change and habitat destruction and fragmentation threatens many plant populations and even entire communities in Mediterranean ecosystems. The Iberian pear, *Pyrus bourgaeana* Decne, a characteristic species of Mediterranean ecosystems, is threatened by both habitat and climate changes. We ask whether and how the growth of mature *P. bourgaeana* in the thermo-Mediterranean zone (i.e., altitude <700 m) has been affected by long-term climate changes during the last century in a fragmented landscape. Dendrochronological

methods were used to find growth–climate relationships. We made the first dendroclimatological analyses and constructed a first 103-year tree-ring chronology (1905–2007) of this species. The tree-ring series revealed large growth variability. We found a clear, strong relationship between tree growth and climate, with annual precipitation being the most important climate factor enhancing radial growth. Our results also showed that warm autumns and winters positively affect growth. There was no temporal stability in the relationship between tree growth and climate. The most general trend was in the relationship between annual precipitation and tree growth: the decrease of rainfall in the last decades of the twentieth century was associated with a constant increase of the correlation coefficient. Water accumulated in the soil in autumn and winter proved to be a key factor augmenting tree growth in the following vegetation period. The climate–growth relationship in *P. bourgaeana* has strengthened in recent decades apparently due to decreased precipitation levels.

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Introduction

Most models of climate change predict global warming associated with a marked variation of the level and spatio-temporal distribution of precipitation and temperature (IPCC 2014; Fischer and Schär 2010; Hoerling et al. 2012). Ecosystems of all climate zones are vulnerable to climate change, even in environments where water is not considered a limited resource (Allen et al. 2010), but according to climate models, the Mediterranean Basin is one of the

world's climate change hotspots because of increases in the frequency, duration, and/or severity of droughts and heat stress (IPCC 2014; Luterbacher et al. 2012; Anderegg et al. 2013; Barbeta et al. 2013). Moreover, land-use change and habitat fragmentation are also important factors rapidly altering the composition of communities in Mediterranean ecosystems (Linares et al. 2010; Olano et al. 2012; Mate-sanz et al. 2015). Therefore, a full understanding of the level of tolerance to climate change is particularly needed in such humanized Mediterranean habitats to forecast population trends and distribution of threatened tree species.

Changes in temperature and water availability often have marked effects on tree growth (Schweingruber 1996). Because trees are long-lived organisms, their ring series record long-term year-to-year changes in climate conditions (Schweingruber 1996; Nicault et al. 2008; Olano et al. 2008, 2012). Due to their complex physiology, trees may respond to climatic changes in complex ways (Drew et al. 2013; Olano et al. 2014; Zang et al. 2014). Moreover, changes in the constraining factors may cause temporal instability in the relationship between climate and tree growth that modulates the response patterns of tree growth (Briffa et al. 1998, 2002; D'Arrigo et al. 2008; Briffa and Matthews 2002; Leburgeois et al. 2012).

In the thermo-Mediterranean zone (i.e., altitude below 700 m; Olson et al. 2001) climate imposes a double adversity for growth and survival of shrubs and trees, i.e., extreme summer droughts and erratic, unpredictable rain (Valladares et al. 2014). Global climate change is generally resulting in higher temperatures and lower precipitation, though such changes are not distributed uniformly throughout the year. As a consequence, the effect of climate change on tree-ring formation can vary depending on the relative importance of the two main limiting climate factors, making predictions particularly difficult. At higher altitudes, e.g., meso-, supra-, and oro-Mediterranean zones (700–2700 m; Olson et al. 2001), where winter cold can be an important stressor, no agreed-upon trend in the response of tree growth to climate change has been inferred from long tree-ring series (Granda et al. 2013, 2014). Some studies have documented an increase (Martinez-Vilalta et al. 2008; Vila et al. 2008; Gimeno et al. 2012; Tegel et al. 2014), and others, a decrease in the growth of Mediterranean trees during the second half of the twentieth century (Jump et al. 2006; Saris et al. 2007; Piovesan et al. 2008; Di Filippo et al. 2010). Changes in climate–growth relationships have also been reported for several Mediterranean tree species (Andreu et al. 2007; Planells et al. 2009; Carrer et al. 2010; Leburgeois et al. 2012). However, the existing data are based mostly on widespread evergreen trees (e.g. Viera et al. 2009; Campelo et al. 2010; Martin-Benito et al. 2011, 2013; Candel-Perez et al. 2012; Gimeno

et al. 2012), and deciduous species have been much less studied in the Mediterranean area (e.g., Tegel et al. 2014; González-González et al. 2015). Even less is known about the response to climate change of small trees and shrubs at low altitudes in the thermo-Mediterranean.

In this study, we examined long-term changes in the stability of the climate–growth relationship of a thermo-Mediterranean deciduous small tree, to gain a perspective on the future persistence of this species, which has been decimated by different components of global change. Our model plant, the Iberian pear (*Pyrus bourgaeana* Decne), is a characteristic species of Mediterranean ecosystems and is considered important to their functioning (Cabezudo and Pérez Latorre 2004; Fedriani et al. 2010; Arenas-Castro et al. 2013). In Doñana National Park (SW Spain), *P. bourgaeana* trees often occur at low density in small scrubland patches among farmland, towns, and other built-over areas (Fedriani et al. 2010). Currently, *P. bourgaeana* shows a very low level of natural regeneration, most likely because of dispersal limitation, extreme summer droughts, and increased herbivore pressure, limiting seedling establishment (Fedriani et al. 2010, 2015). It is unclear whether and how mature *P. bourgaeana* tree growth has been affected by long-term changes in climate through the last century. This should be a fundamental step to predict the fate of the species, and, in general, of other deciduous thermo- and meso-Mediterranean shrubs and small trees, under a scenario of global warming. Specifically, in this study, we sought to answer the following two questions: (1) what sort of long-term climate–growth relationship is shown by *P. bourgaeana*?, and (2) is this relationship temporally stable, and if not, what is the trend of such/these changes, (if found any)?

Methods

Study site

The study site is located in southwestern Spain, in the Doñana World Biosphere Reserve (37°9'N, 6°26'W; 0–80 m a.s.l.) on the west bank of the Guadalquivir River estuary. The Doñana area comprises three main ecosystems: mobile dunes, scrubland (where *P. bourgaeana* grows), and marshes. In the tree layer, in addition to our focal species, there are scattered *Quercus suber*, *Olea europaea* var. *sylvestris*, *Fraxinus angustifolia*, and *Pinus pinea*. In the understory, there are *Pistacia lentiscus*, *Halimium halimifolium*, *Ulex* spp., *Chamaerops humilis*, and *Erica* sp. (Valverde 1958; Fedriani et al. 1998, 2010). *P. bourgaeana* trees occur at low densities (generally less than one individual per ha; Fedriani et al. 2010) in patches of Mediterranean scrubland that are isolated from each other by natural or anthropogenic barriers (marshes, sand

dunes or cultivations) (see also Żywiec et al. 2012 for details).

Climate data

Climate data were obtained from the Spanish Agencia Estatal de Meteorología (AEMET) meteorological station located in Sevilla, the closest (ca. 80 km away) meteorological station providing long-term data and located near sea level (11 m a.s.l.). Continuous data for daily mean temperature and daily total precipitation were available since 1946. We calculated mean temperature and accumulated precipitation for each month, as well as annual accumulated precipitation. Annual accumulated precipitation was summed from the September of the previous year to the August of the current year.

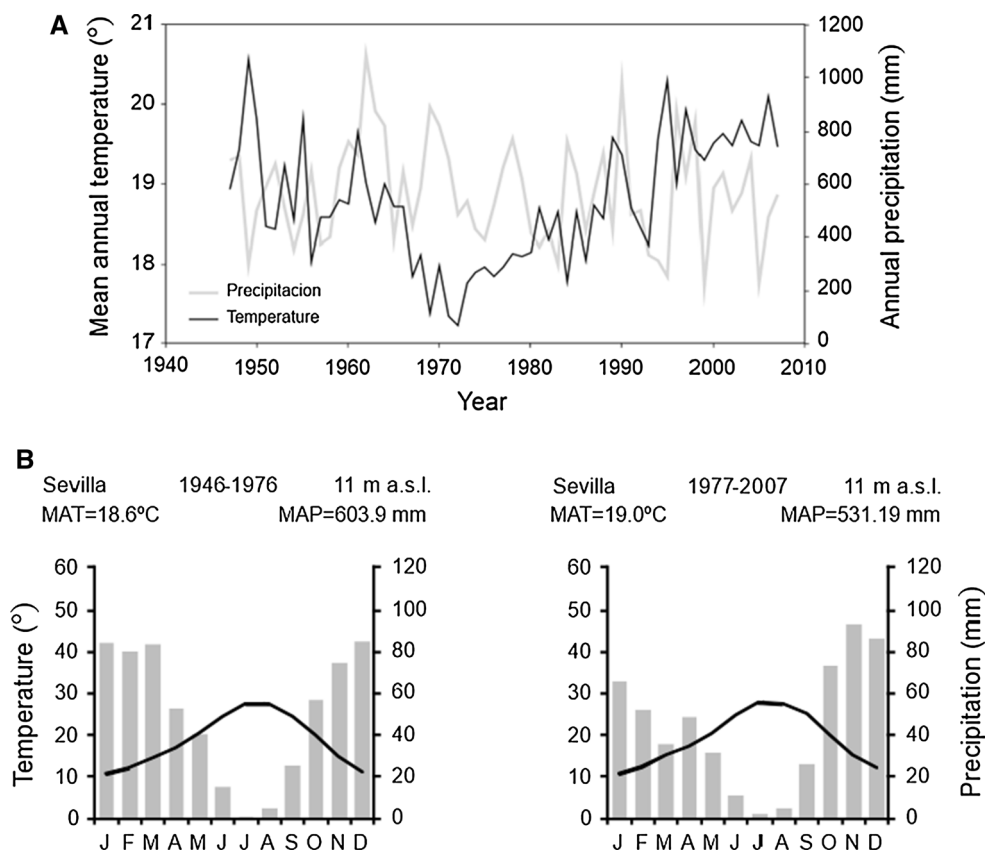
The climate is Mediterranean sub-humid, and is characterized by dry, hot, long summers (June–September) and mild, wet winters (November–February). In the study period (from 1946 to 2007), mean annual temperature varied considerably (Fig. 1a). The highest temperatures were recorded at the beginning and at the end of our studied period. The lowest temperatures occurred in the early 1970s and then increased consistently until the first decade of the twenty-first century. Total annual precipitation also varied markedly, but no conspicuous temporal trend was found (Fig. 1a;

$R^2 = 0.003$, $F_{1,59} = 1.20$, $P = 0.28$). Groundwater availability shows marked seasonal changes in this area (Instituto Tecnológico Geominero Español, 1992). Precipitation is highest in the winter and lowest in the summer. The distribution of precipitation in a year changed along a studied period. The precipitation of January, February and March decreased considerably and the precipitation of October, November and December increased in the second half of the period considered in the study (Fig. 1b).

Study species

Pyrus bourgaeana Decne (*Rosaceae*) is a broadleaved deciduous tree. It is native to the Iberian Peninsula (Spain, Portugal) and North Africa (Morocco) (Aldasoro et al. 1996). It typically reaches 3–6 m height. Budburst occurs from February to April, and leaf senescence from June to September (Mediavilla and Escudero 2003; Arenas-Castro 2012). It flowers during February–March, is hermaphrodite and self-incompatible (Żywiec et al. 2012; Authors, unpublished data). In Doñana, annually, each *P. bourgaeana* tree typically produces between 200 and 450 fleshy fruits which ripen during the autumn (September–November; Fedriani et al. 2015). The fruits are non-dehiscent globose pomes (2–3 cm diameter) weighing ca. 6.7 g, with sugary, juicy pulp (Fedriani et al. 2012).

Fig. 1 a Mean annual temperature (1946–2007) and total annual precipitation (1946–2007) near the study area; data from Sevilla weather station and **b** Climate diagrams for the meteorological station in Sevilla (37°23'10"N, 5°59'33"W, 11 m a. s. l.) for the periods 1946–1976 and 1977–2007 indicating the intra-annual distribution of monthly mean temperatures (*line*) and monthly total precipitation (*bars*); MAT mean annual temperature (°C) and MAP mean annual precipitation [mm]



The Doñana population of *P. bourgaeana* has limited reproduction and regeneration ability (Fedriani and Delibes 2009; Fedriani et al. 2010, 2012). Our unpublished data indicate that the Doñana population of this long-lived tree shows a marked left-skewed demographic structure, with many individuals in older age classes, few juveniles, and even fewer seedlings and saplings. Despite the sparseness of this population, *P. bourgaeana* fruit, seeds, seedlings, and leaves represent important resources for diverse animal guilds, especially during the dry summers (e.g., Fedriani and Delibes 2009; Fedriani et al. 2012); so, its persistence is critical to the ecological stability of the community.

Data collection

Trees ($n = 32$) were sampled in winter and spring of 2008 and 2009. Tree clusters grow scattered in Mediterranean scrubland, in similar light conditions representing the uppermost vegetation layer. Their mean diameter (dbh) was 18 cm (± 11). The trees were cored with an increment borer (diameter 5 mm) ca. 40 cm above the tree base. The cores were glued to wood slats, dried and polished with a belt sander (progressively up to grid 600) to make the tree-ring sequences clearly visible. The samples were scanned at 2400 dpi resolution. WIN-DENDRO (http://www.regentinstruments.com/assets/windendro_about.html) was used to measure the annual tree-ring widths to 0.01 mm accuracy. Core sections with extremely narrow, missing or false rings were carefully checked under a stereomicroscope at 7–45 \times . The quality of dating was checked with the COFECHA program (Dendrochronology Program Library, Holmes and Cook 1983). All measurement series with potential errors were re-checked once again and corrected if dating errors were found (Holmes et al. 1986).

Data analysis

Tree-ring chronology

We calculated the index series to study the influence of climate factors on tree growth (Holmes 1994). Individual raw ring-width series were standardized using ARSTAN software by fitting a horizontal line through the mean values of the series, and, in the next step, by dividing the measurement in every year by the mean value calculated for each series (Holmes 1994). Then, we performed autoregressive modeling of the index series to remove autocorrelations and enhance the common signal of the trees. After that, the indices of individual trees were averaged to an indexed chronology by applying a biweight robust mean (Holmes 1994). The first-order autocorrelation of the 1904–2007 raw chronology (a measure of the lagged effect of ring width from the previous year on next-season

growth) was 0.455, and the first-order autocorrelation of the indexed chronology was -0.137 . This means that a great part of the autocorrelation was removed in the process of constructing the indexed chronology allowing dendroclimatological analyses.

Mean sensitivity, expressing the interannual variability of ring width, was calculated for the indexed chronology 1947–2007 (Fritts 1976). For the period 1947–2007, we also calculated the expressed population signal (EPS), which quantifies how well a chronology based on a finite number of trees represents the hypothetical perfect or true chronology (Wigley et al. 1984).

Weather factors influencing tree growth and temporal changes in the climate–growth relationship

The indexed chronology was used for calculation of the tree growth–climate relationship. To evaluate the potential relationship between tree growth and annual precipitation, we performed several simple regression analyses, first regressing the indexed chronology and year, and then, annual precipitation and year. Then, we used the residuals from those two previous analyses and regressed them to evaluate their potential relationship once the effect of year was corrected for (e.g., Ovaskainen et al. 2013).

The relationships between tree growth and monthly climatic conditions (mean monthly temperature, total monthly precipitation) were calculated through correlation and response function analyses with DendroClim2002 (Biondi and Waikul 2004). The program applies a bootstrap process to assess the statistical significance of the correlations.

Time-dependent changes in the relationship between tree growth and climate factors were analyzed using DendroClim2002 as well. The correlations of indexed chronology with total monthly precipitation and mean monthly temperature (from previous September to current June) were done for 30-year periods. The moving correlation windows (32 periods) were progressively shifted by 1 year over time at each iteration (the first for 1947–1976 and the last for 1978–2007). Moreover, for each 30-year period, climatic conditions were defined as mean monthly precipitation and mean monthly temperature.

Results

Patterns of tree growth

The constructed indexed chronology (based on 32 *P. bourgaeana* trees) covers a 103-year time span (1905–2007; Fig. 2a, b). The two oldest trees had 104 rings (at ca. 40 cm height). Mean ring width was 1.34 mm

(± 1.04) for 1947–2007, and 1.35 mm (± 1.04) for 1904–2007; it reached maximum of 7.83 mm in 1996 as calculated from the raw chronology. The mean sensitivity of the indexed chronology for 1947–2007 was 0.49. The expressed population signal (EPS) for 1947–2007 was 0.87. With time, the number of years for which the tree series showed no ring formation (missing ring) increased: in the 1950s and 1960s, there were no absent rings; in the 1970s, 1 year had no ring (1974; 2 trees); in the 1980s, there were 4 such years (1980, 1981, 1983, 1989; 1 tree each year); in the 1990s, there were 5 such years (1992–1995 and 1999; 1–3 trees each year); and in the first 7 years of the twenty-first century, there was 1 year with a missing ring (2005; 1 tree).

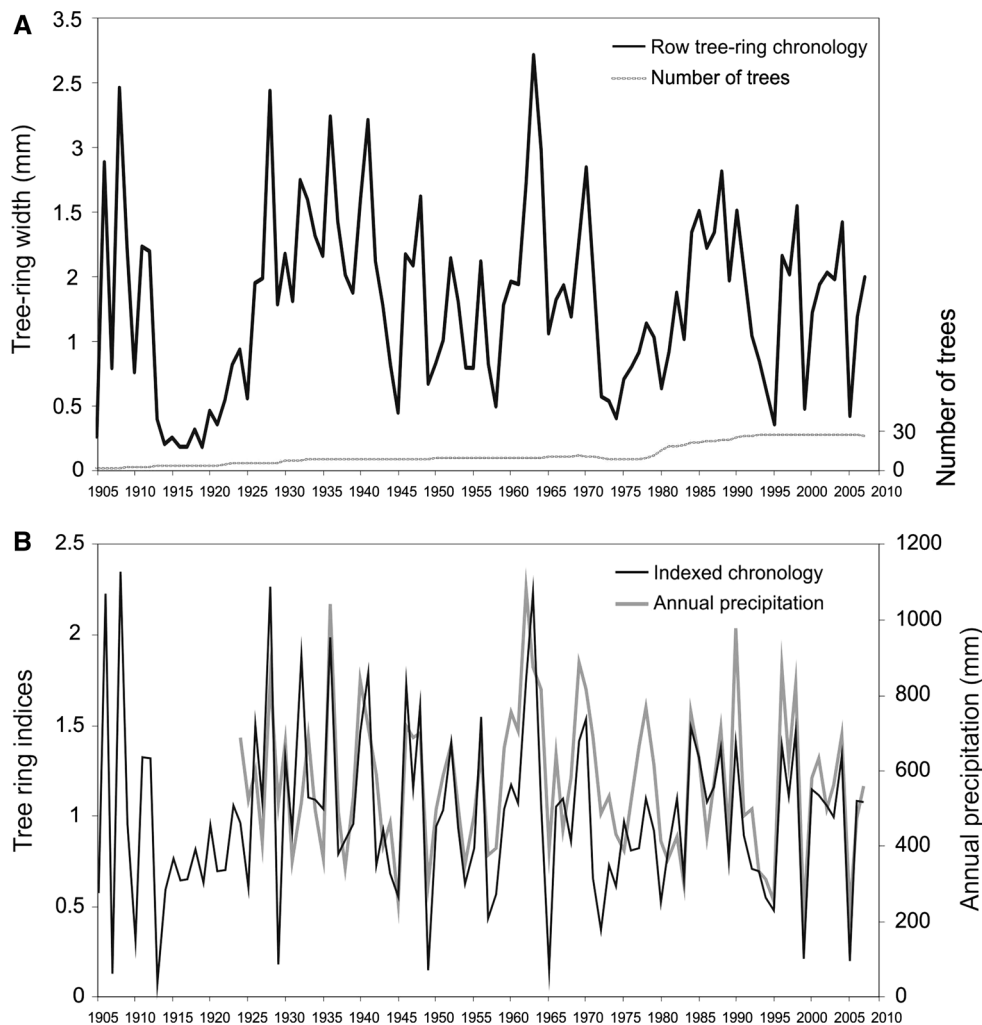
Long-term climate–growth relationship

Tree growth was strongly and positively related to annual precipitation, summed from September of the previous year to August of the current year (Fig. 2b; regression of

residuals: $R^2 = 0.66$, $F_{1,59} = 114.25$, $P < 10^{-4}$). Tree growth was not correlated with total precipitation of the preceding year (regression of residuals: $R^2 = 0.01$, $F_{1,58} = 0.28$, $P = 0.60$). Correlation coefficients showed that precipitation from November of the previous year to January of the current year was especially important for tree growth (Fig. 3). Almost the same was found using response function coefficients; precipitation of previous October, previous November, and January and February of the current year positively affected growth (Fig. 3). Precipitation from February to June of the current year also positively influenced growth, with June precipitation being the most important among those months (Fig. 3).

Mean monthly temperature influenced tree growth less than precipitation did (Fig. 3). There was a positive correlation between tree growth and temperature of the previous October. High winter temperature (previous December and current January) also influenced growth positively, but this relationship was significant only for current January temperature. High temperature from April

Fig. 2 Row chronology of *P. bourgaeana* in the Doñana area with number of measured trees (a) and indexed chronology with annual precipitation records from the Sevilla weather station (b)



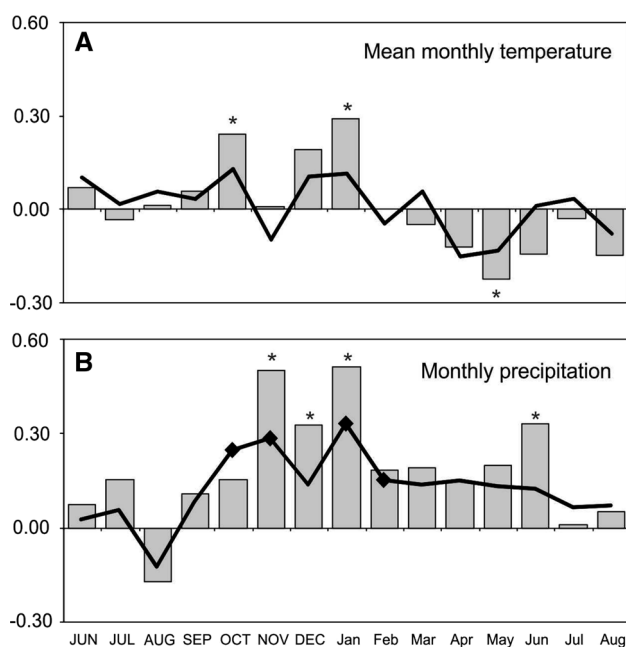


Fig. 3 Bootstrap correlations (*bars*) and response (*line*) calculated between the indexed chronology of *P. bourgaeana* and monthly mean temperature (**a**) and sum of monthly precipitation (**b**) from previous June (JUN) to current August (Aug) in 1947–2007; *asterisks* (correlation) and *marks* (response) indicate significance at 95 % confidence level

to June of the current year negatively influenced tree-ring formation, but this relationship was significant only for May (Fig. 3). Response function coefficients between monthly temperatures and growth were not significant in any month, although they showed similar trends as for the correlation functions (Fig. 3).

Changes in the climate–growth relationship

The relationship between tree growth and total annual precipitation was always positive and strong during the whole analyzed period (1947–2007; Fig. 4). The influence of annual precipitation becomes more important for tree growth in the late twentieth and early twenty-first centuries, when annual precipitation decreased and it probably became most limiting. For total monthly precipitation, during the whole analyzed period, there was a strong positive relationship between tree growth and precipitation in the previous November and current January (Fig. 5). The relationship between total monthly precipitation of the previous December and tree growth strengthened beginning in the 1960s (Fig. 5). The relationship between precipitation in the spring and tree growth fluctuated and was rather weak in the studied period (Fig. 5). The positive relationship between growth and total precipitation in the vegetative period was strongest for June, but it decreased in the late twentieth and early twenty-first centuries (Fig. 5).

The positive relationship between tree growth and mean temperature in autumn and winter (from previous October to December) strengthened in the second half of the studied period, coinciding with an increasing mean temperature trend; we found the same trend for February (Fig. 5). There was no time trend in the relationship between tree growth and mean spring temperature, despite the increase of temperature in the last few decades (Fig. 5).

Discussion

In this study, we made the first dendroclimatological analyses of the thermo-Mediterranean tree *P. bourgaeana* and constructed a first 103-year tree-ring chronology of this species. Recently, Arenas-Castro et al. (2015) described average increase in basal area per year and age structure based on *P. bourgaeana* tree-rings. This small deciduous tree grows in the thermo-Mediterranean zone, where winter is relatively mild, and thus, tree responses to climate change could deviate from those observed at higher Mediterranean altitudes. We found a strong relationship between tree growth and climate, with annual precipitation being the most important limiting factor for *P. bourgaeana* radial growth. We showed that in recent decades, the relationship between growth and climate has strengthened. The increasing number of years with absent rings suggests that climatic conditions for this species worsened toward the end of the twentieth century. In combination with habitat destruction and fragmentation in the Doñana area, and in the Mediterranean in general, it indicates that our study species is seriously threatened by a myriad of contrasting factors.

Tree growth and its long-term relationship with climate

The high sensitivity of *P. bourgaeana*, as expressed in large year-to-year variation of ring width, might have been partly related to the species's low stomatal sensitivity to atmospheric drought (Mediavilla and Escudero 2003). The high sensitivity is typical for trees growing in environments with a strong, temporally variable limiting factor (Holmes et al. 1986). *P. bourgaeana* mean sensitivity (0.49) was close to values found for the small-sized *Pinus edulis* (0.58) growing at extremely dry sites in Arizona (Fritts et al. 1965), but much higher than most values reported recently for large-sized tree species growing in the Mediterranean area at higher altitudes, such as *Pinus sylvestris*, *P. nigra*, *Abies alba*, and *Fagus sylvatica* (range 0.19–0.35; Leburgeois et al. 2012; Martin-Benito et al. 2013; Andreu et al. 2007; Cailleret and Davi 2011; Piovesan et al. 2008).

Fig. 4 Bootstrap correlations (black line) of *P. bourgaeana* indexed chronology with total annual precipitation. Bootstrap correlations were calculated for 30-year periods lagged year-by-year through 1947–2007 (from 1947–1976 to 1978–2007). Grey dashed line is trend of mean annual precipitation over the 32 30-year periods

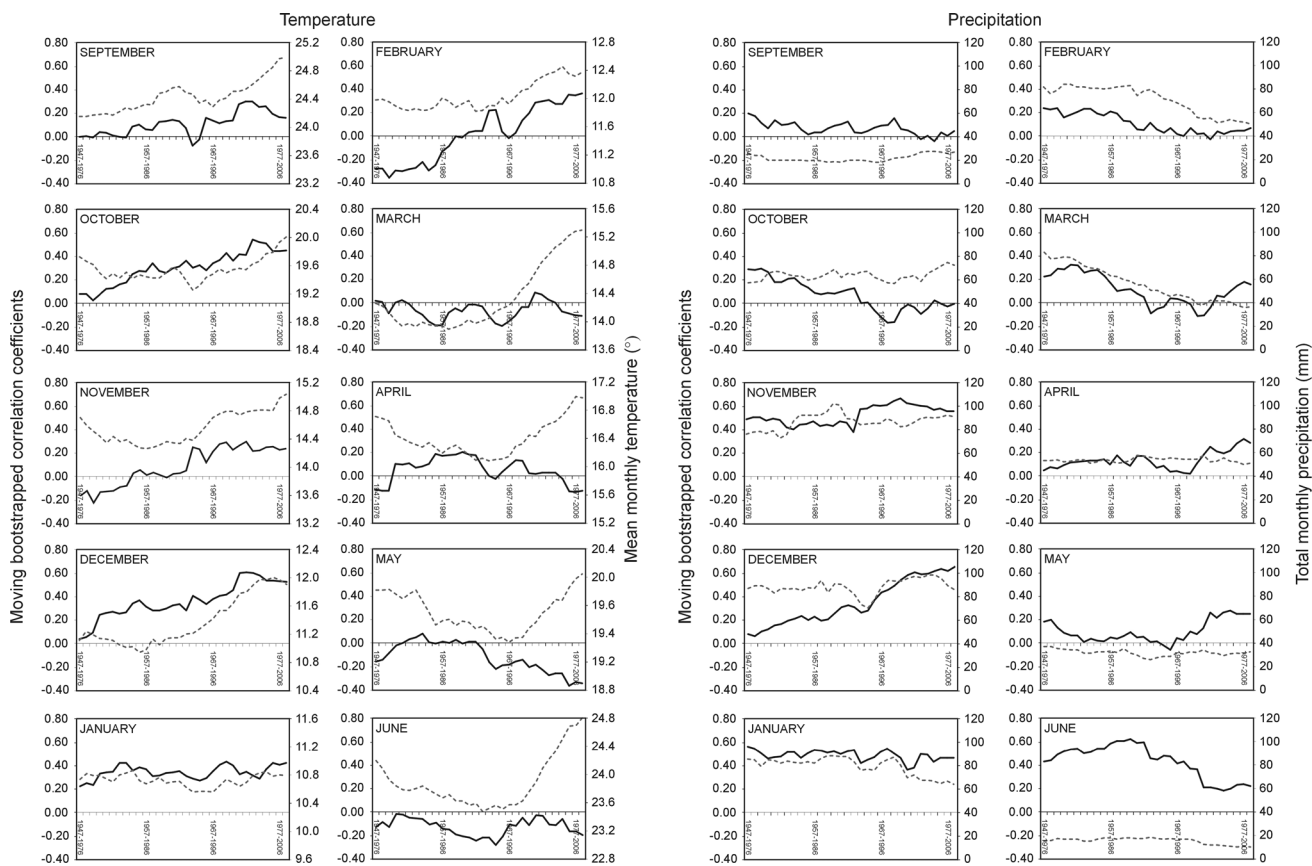
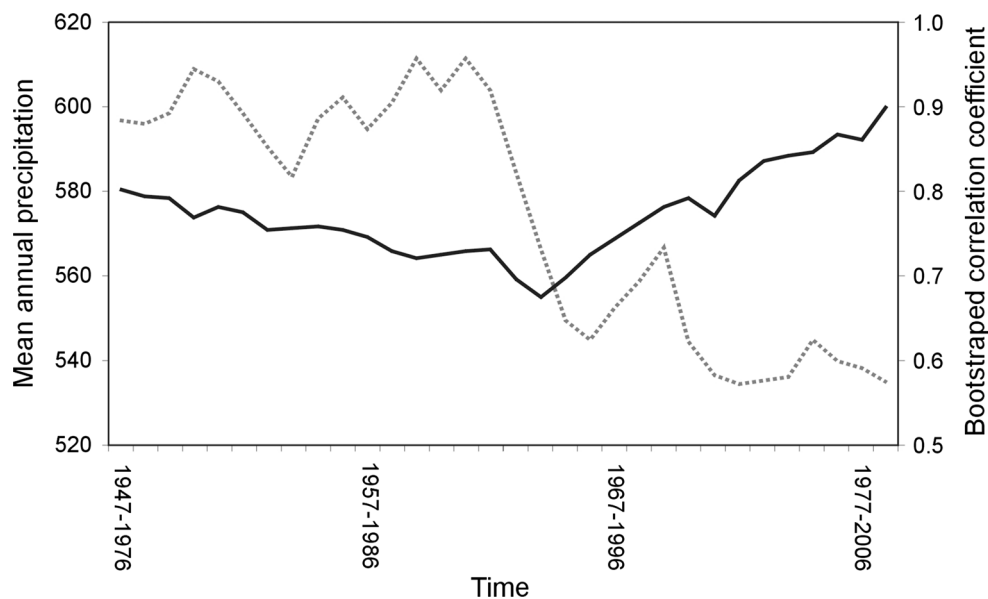


Fig. 5 Moving bootstrap correlations (black line) of *P. bourgaeana* indexed chronology with total monthly precipitation and mean monthly temperature (from previous September to current June) in 1947–2007. Bootstrap correlations were calculated for 30-year

periods lagged year-by-year through 1947–2007 (from 1947–1976 to 1978–2007). Grey dashed line is trend of total monthly precipitation and mean monthly temperature over the 32 30-year periods

Drought is a primary growth-limiting factor of Mediterranean tree species, especially at low altitudes (Thompson 2005; Vicente-Serrano et al. 2010, 2012;

Pugnaire and Valladares 2007; Olano et al. 2014). However, previous dendroclimatological studies of a number of tree and shrub species at relatively high altitudes show that

tree-ring chronologies are not always strongly correlated with annual precipitation (Cherubini et al. 2003; Gimeno et al. 2012; but see Gutiérrez et al. 2011; Martin-Benito et al. 2011; Granda et al. 2013). In our thermo-Mediterranean study, *P. bourgaeana* showed a very strong positive relationship between total annual precipitation and radial growth.

Previous autumn and winter precipitation is known to boost tree-ring formation in several Mediterranean species (e.g., Campelo et al. 2006, 2010; Di Filippo et al. 2010). The highest precipitation in this region occurs in winter or autumn, and this time is responsible for soil water recharge (Instituto Tecnológico Geominero Español 1992). However, maximum tree growth occurs in the spring (Cherubini et al. 2003). In the study area, precipitation was relatively high from November to January, and its year-to-year variation was lower in these months than during the rest of the year. Precipitation in the spring months of the growth season was much more variable, and trended downward over the course of the studied period. In the Doñana area, water accumulated in the ground during autumn and winter, when *P. bourgaeana* trees are dormant, seems to be a key factor influencing their growth in the following vegetative period. It establishes an essential resource for initiation of cambial activity and production of early wood (Granda et al. 2013; Martin-Benito et al. 2013).

Temperature limits tree growth in some Mediterranean ecosystems, though less than precipitation does (Cherubini et al. 2003). The limiting effect of winter cold has been observed mainly for evergreen species (Martin-Benito et al. 2011, 2013; Granda et al. 2013) and for trees growing in mountain conditions (Leburgeois et al. 2012; Martin-Benito et al. 2013). Our results in a warmer thermo-Mediterranean site show that a warm previous autumn and winter positively affect growth of the deciduous species growing at low elevation. Two mechanisms might explain such a relationship in *P. bourgaeana*. First, high autumn temperatures can increase the quality, size and survival of buds formed for the next year, and, thus, the assimilation ability of the plant in the following season (Weber et al. 2007). Second, a warm end of winter can accelerate cambial reactivation after winter dormancy and earlier leaf burst, thereby prolonging the growing season and the period of wood formation (Leburgeois et al. 2012; Rötzer et al. 2004; Sanz-Pérez et al. 2009; Cailleret and Davi 2011; Viera et al. 2014). On the other hand, higher temperatures in spring and early summer slightly reduced the growth of *P. bourgaeana*. In water-limited regions, high temperatures can cause water stress and the trees usually respond by reducing their stomatal conductance to avoid water loss (Baquedano and Castillo 2007) or, in the case of *P. bourgaeana*, by early leaf abscission in the summer (Mediavilla and Escudero 2003).

Changes in the climate–growth relationship

Climatic changes have led to temporal instability of climate–growth relationships in the Mediterranean region (Andreu et al. 2007; Planells et al. 2009; Carrer et al. 2010) and elsewhere (Briffa et al. 1998). Generally, our results showed no temporal stability in the relationship between tree growth and climate from 1947 to 2007. The most general temporal trend was the relationship between annual precipitation and tree growth. The decrease of rainfall in the last decades of the twentieth century caused a steady rise of that correlation, reaching high values. This relationship was very strong during the whole period studied, but its strengthening through time makes it clear that lower precipitation during the last decades has limited the growth of *P. bourgaeana* more than it did earlier. Similar long-term patterns have been reported from other species living at higher altitudes, e.g., *Fagus sylvatica* (Piovesan et al. 2008) and for *Pinus sylvestris* (Planells et al. 2009; Andreu-Hayles et al. 2011).

The distribution of rainfall over the year is the most essential element of the precipitation–growth relationship. The most stable and strongest relationship was found between growth and total monthly precipitation in the previous November and January. Spring drought is considered a primary growth-limiting factor in the high-altitude Mediterranean areas (Leburgeois et al. 2012; Cherubini et al. 2003). Surprisingly, during some spring months, the relationship between *P. bourgaeana* growth and precipitation was stronger at the beginning of the studied period, when mean precipitation was higher, and almost disappeared in the late twentieth and early twenty-first centuries, when precipitation was lower. This suggests that high precipitation in the spring boosts *P. bourgaeana* growth, but that lower precipitation at that time does not limit growth so much, as it depends mostly on autumn and winter precipitation. On the other hand, the relationship between June precipitation and growth decreased strongly over time even though June precipitation was very low throughout the studied period what can be the result of the changed tree phenology.

In the twentieth century, directional changes in the temperature–growth relationship have been reported by other researchers (Andreu et al. 2007; Planells et al. 2009). Our data show that the positive relationship between autumn and winter temperature and tree growth became stronger at the end of the twentieth century, coincident with the increasing trend of temperature in this period. It seems that winter warming initiated a positive growth response to temperature. Leburgeois et al. (2010, 2012) found that spring warming strengthened the positive growth response to temperature in several species. In our study, the positive

relationship between mean February temperature and tree growth started when temperature during this month increased. This suggests that at higher February temperature, the growth of *P. bourgaeana* begins earlier, extending the vegetative period.

Our investigation indicates that the growth of mature *P. bourgaeana* in Doñana was limited by climatic conditions during the whole studied period, and that the climate–growth relationship was more pronounced in the late twentieth and early twenty-first centuries as climate became more arid. There was a clearly stronger relationship between precipitation and tree growth at the end of the twentieth century, when the number of missing rings also increased. However, the negative influence of climate change was not as strong as might be expected. A recent review of results on current effects of climate change on Mediterranean forests reached a similar conclusion (Valladares et al. 2014). As for any other tree species, however, the persistence of *P. bourgaeana* in Doñana depends not only on the vigor and health of existing mature trees, but also on recruitment (Fedriani et al. 2012). Though adult trees might somewhat tolerate climate change, the current scarcity of young *P. bourgaeana* plants (seedlings, saplings; our unpublished data) appears to pose a serious threat to this population. Further similar analyses of other thermo-Mediterranean tree and shrub species growing in such particularly stressful conditions are needed to achieve a better understanding of the response of plants to climate change in humanized landscapes.

Author contribution statement M.Ž., J.M.F. and M.D. conceived the ideas; M.Ž., E.M., T.Z. and G.C. collected the data; E.M., M.Ž. and T.Z. analysed the data; and M.Z., E.M. and J.M.F. led the writing.

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Compliance with ethical standards

Conflict of interest The authors declare that they have no conflict of interest.

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